U-Pb zircon geochronology and Sr-Nd isotopic characteristic of Late Neoproterozoic Bornaward granitoids (Taknar zone exotic block), Iran

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Abstract: The study area (Bornaward granite) is located in northeast of Iran (Khurasan Razavi province), about 280 km southwest of Mashhad city. Taknar zone is an exotic block, bordered by two major faults, Great Kavir fault (Drouneh) to the south and Rivash fault in the north. A complex of granite, granodiorite, monzonite and diorite crop out at the center of Taknar zone. They are named as “Bornaward granite”. Published data using Rb–Sr whole-rock and biotite isotopic methods on granitoid rocks (Bornaward granite) gave ages of 154 to 111 Ma. The results of U-Pb zircon dating of granodiorite is 552.69 ± 10.89 Ma and granite is 538.22 – 1.82, + 4.28 Ma (Late Neoproterozoic time). Both granite and granodiorite are classified as belonging to the ilmenite-series of reduced S-type granitoids. Chemically, they are per-aluminous, high-K calc-alkaline with relatively enriched in LILE, Rb, K and depleted in Sr, Ba, Nb, Ti, Ta, Y and Yb. Chondrite-normalized Rare Earth Element (REE) plots indicate minor enrichments of light REE in composition with heavy REE, with (La/Yb)N between 3.5-5.6 and high total REE (193-252) with strong negative anomaly of Eu. They have an initial 87Sr/86Sr and age published by Soltani [5] is incorrect. Based on paleontological studies, the age of Taknar Formation is Ordovician [4]. Since the age of granitoid which are intruded Taknar Formation are Precambrian (present study), therefore this age is also incorrect. We present the results of new radiogenic isotope analyses and more accurate U-Pb in zircon age dating and discuss the paleotectonic and petrogenetic implications of these new data for Bornaward Neoproterozoic granitoid.

Keywords: Taknar; Bornaward; Rb-Nd; U-Pb-zircon; late Neoproterozoic.

Introduction
The study area (Taknar mine area) is located in northeast of Iran (Khurasan Razavi province), about 280 km southwest of Mashhad city and 28 km northwest of Bardaskan (Fig. 1). Taknar zone consists of Precambrian, Paleozoic, Mesozoic and Tertiary rocks [1-3]. Taknar zone is an exotic block bordered by two major faults, Great Kavir fault (Drouneh) to the south and Rivash fault in the north (Fig. 1).

Taknar Formation covers a large area within the Taknar zone. Based on the previous studies, age of Taknar Formation is Precambrian [1]. However, recent Palynological studies show that the age of Taknar Formation is Ordovician [4].

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We found that the initial 87Sr/86Sr and age published by Soltani [5] is incorrect. Based on paleontological studies, the age of Taknar Formation is Ordovician [4]. Since the age of granitoid which are intruded Taknar Formation are Precambrian (present study), therefore this age is also incorrect. We present the results of new radiogenic isotope analyses and more accurate U-Pb in zircon age dating and discuss the paleotectonic and petrogenetic implications of these new data for Bornaward Neoproterozoic granitoid.
Geology

Taknar Formation is divided into lower, middle and upper members [1]. The lower member crops out in the central part of Taknar with thickness of 1500m. The lower member consists of meta-tuff and meta rhyolite. The middle member is exposed along the northern and southern margins of the Taknar zone [1]. The thickness of middle member varies between 150 to 350m [1]. It consists of alternation of meta–carbonate and quartzite with minor meta-rhyolite. The Upper member consists of fine to coarse-grained meta–graywakes (slate to phylite). Between Dahan Qaleh and Drouneh, the upper Taknar Formation is followed by Infracambrian and Early paleozoic dolomite and quartzite [1].

A complex of intrusive rocks crop out at the eithen of Taknar zone (Fig. 2). They are named as “Bornaward granite” [3, 6, 7]. The intrusive rocks are composed of granite, granodiorite, monzonite and diorite [8]. Granite and granodiorite are dominant. Based on field observation, diorite is older than granite and granodiorite [8]. Several episodes of magmatism were recognized at Taknar (Bornaward granitoid) [8]. Soltani [5] did some Rb-Sr analyses and calculated the age of intrusive rocks between 154 to 111 Ma. This age is not correct (See below). Petrography of intrusive and metamorphic rocks was studied by Homam [9] and Sepahigrou [10]. Zirjanizadeh [11] worked on petrography, major oxides and trace elements of intrusive rocks. Detail field work, petrography, Petrogenesis, radio isotopes and U-Pb zircon age are carried by the author.

Taknar mine is a magnetite-rich polymetal (Cu, Zn, Au, Ag, and Pb) massive sulfide deposit [12]. Mineralization shows good layering, therefore, it is a syngenetic deposit type and was formed at certain horizon within the upper member of Taknar Formation [12, 13]. Malekzadeh and Ghoorch worked on the geology, alteration and mineralization of Taknar deposit [12, 13]. Based on local name, from north to south, deposits are named as Tak-II, Tak-I, Tak-IV and Tak-III.

Several episodes of metamorphism have affected the Taknar Formation and the intrusive rocks.
1- In areas where the polymetal massive sulfides are formed, the rocks are altered due to hydrothermal fluid; different types of alteration zones such as chlorite, sericite-chlorite, chlorite-carbonate and silicification were formed [12].
2- Contact metamorphism was developed at the contact of intrusive rocks with Taknar Formation.
3- Taknar Formation and most of the intrusive rocks were subjected to low-grade regional metamorphism. Slate and sericite-chlorite schist were formed within some parts of the Taknar Formation (areas with tuff, shale and graywacke). Regional metamorphism may have occurred at different episodes. 1- Precambrian during and after the granitoids intrusion; 2- Late Paleozoic; and 3- Middle Jurassic.

Figure 1. Taknar zone is shown on Landsat image.
Taknar zone is an exotic block and it is located between two major faults, therefore since Precambrian, major displacement took place. Both right and left-lateral strike-slip faults are observed in the study area (Fig 2, 3). At least three generations of strike-slip faults are observed trending: 1- NE-SW, 2- NW-SE and 3- N-S.

Analytical Techniques

Bulk-rock chemistry

Major elements were analyzed by wavelength-dispersion X-ray fluorescence spectrometry (XRF) using fused discs. Pressed powder pellets were used for Zr, Nb, Sn, Sr, Ba, Sc, Y and Ga measurements by XRF (The XRF spectrometer used in this study was a Philips PW 1410) at Ferdowsi University of Mashhad, Iran. REE and trace elements composition of rocks was determined by ICP-MS in Canada (Acme Lab. Canada).

U/Pb dating

Two rock samples, which were both analyzed for Sr and Nd isotopes, were chosen for zircon U-Pb age dating. Zircons were isolated using standard mineral separation techniques. From each rock about 70 zircon grains were separated. Zircons were mounted along with a zircon standard and a couple of chips of NBS 610 Trace Element Glass in epoxy and polished down to 20 μm. Zircon age dating was done at the Arizona LaserChron Center with the method of Gehrels, and Valencia, (2006). Cathode-luminescence (CL) images are acquired for samples to be analyzed because they provide a powerful tool for placing laser pits in homogeneous portions of crystals, and also can help determine the origin (e.g., igneous, metamorphic, hydrothermal) of zircon grains. Laser ablation takes place with a beam diameter of either 35 or 25 microns for most applications, or with a beam diameter of 15 or 10 microns if finer spatial resolution is needed. With a 35 or 25 micron beam, the laser is set at a repetition rate of 8 Hz and energy of 100 mJ, which excavates at a rate of ~1 micron per second. This generates a signal of ~100,000 cps per ppm for U in zircon. For smaller beam sizes, the ablation rate is reduced to ~0.5 micron/second by reducing the laser energy (60 mJ) and repetition rate (4 Hz). In both cases the ablated material is removed from the ablation chamber in He carrier gas. The carrier gas (and sample) are then mixed with Ar gas before entering the plasma of the Inductively Coupled Plasma Mass Spectrometer (ICPMS). Isotopic analysis is performed with a Multicollector Inductively Coupled Plasma Mass Spectrometer (GVI Isotope) that is equipped with an S-option interface. The instrument is equipped with a collision cell that is operated with a flow rate of 0.2 ml/min of argon to create a uniform energy distribution, and the accelerating voltage is ~6 kV. Collectors include nine Faraday detectors and four low-side channeltron multipliers, all of which are moveable, as well as an axial Daly photomultiplier.

Nd–Sr isotopes

Sr and Nd isotopic analyses were performed on a 6-collector Finnigan MAT 261 Thermal Ionization Mass Spectrometer at the University of Colorado, Boulder (USA). 87Sr/86Sr ratios were analyzed using four-collector static mode measurements. Thirty measurements of SRM-987 during study period yielded mean 87Sr/86Sr=0.71032 ± 2 (error is the 2 sigma mean). Measured 87Sr/86Sr were corrected to SRM-987= 0.71028. Error in the 2 sigmas of mean refer to last two digits of the 87Sr/86Sr ratio. Measured 143Nd/144Nd normalized to 146Nd/144Nd=0.7219. Analyses were dynamic mode, three-collector measurements. Thirty-three
measurements of the La Jolla Nd standard during the study period yielded a mean $^{143}\text{Nd}/^{144}\text{Nd} = 0.511838 \pm 8$ (error is the 2 sigma mean).

**Petrography**

The granitic rocks of the Bornaward cut through the metamorphic rocks of the Taknar Formation. Based on field investigations and microscopic studies, several intrusive rocks are recognized in the study area (Figs. 3, 4). They are mainly granite, granodiorite and minor diorite-gabbro. Since the Taknar zone is an exotic block and it is surrounded by two major faults (Drouneh and Taknar), several system of faulting are recognized within the Taknar zone. The majority of the faults are strick-slip. One of the right lateral strick-slip fault cut the granitic rocks and the massive sulfide ore body with about 500 meter displacement (Figs. 3, 4). Both regional metamorphism and faulting caused some changes in fabric of the granitic rocks. In some locations, along the fault gneissic texture was formed.

**Granite & biotite-granite:** These rocks crop out in both Tak-I and Tak-II (Figs. 3, 4). It has graphic to granular textures. In Some locations granite is strongly deformed with a NEE– trending foliation. The granite contains 45%–50% quartz, 30%–35% K- feldspar, 5% plagioclase, and 1-5% green biotite. Accessory minerals are zircon, and ilmenite.

**Granodiorite:** This unit also crops out in both Tak-I and Tak II, in Tak-I it cuts massive sulfide ore body and Taknar Formation (Figs. 3, 4). The texture is porphyry with some graphic intergrowth. Granodiorite contains 10% quartz (phenocrysts), 3-7% K- feldspar, 5-8% plagioclase, and 0.5 % biotite. Accessory minerals are zircon, and ilmenite. Quartz as veinlets formed near the contact of granodiorite with country rocks.

**Gabbro-diorite:** It crops out near massive sulfide ore bodies both in Tak-I and Tak II (Figs. 3, 4). Since it doesn’t have zircon, therefore, we could not get the age of this rock based on U-Pb in zircon. Based on field observation and chemical composition (low SiO$_2$, met-aluminous, volcanic arc setting) it seems that the gabbro-diorite is older than granite and granodiorite. Stockwork mineralization (quartz-magnetite-chalcopyrite) is found within the gabbro-diorite both in Tak I and Tak II, but the granodiorite is cutting the mineralized rocks. This is also a good evidence that gabbro-diorite is older than granodiorite-granite. It has porphyry texture. It contains 55% plagioclase and 40% pyroxene (aegirine-augite). Due to regional metamorphism and alteration both plagioclase and pyroxene are altered to chlorite and carbonate.

Spotted granite: It crops out only in Tak-II (Fig. 3). It has porphyry textures with elongated black ellipsoids (Figs. 5, 6). These elongated ellipsoids formed during regional metamorphism and along the fault movements. The ellipsoids are consist of chlorite, quartz & opaque minerals.

![Figure 3. Geological map of Taknar, Tak-II.](image-url)
**Figure 4.** Geological map of Taknar, Tak-I.

**Figure 5.** Field photograph of spotted granite showing elongated ellipsoid.

**Figure 6.** Photomicrograph of spotted granite showing that the spots are consisted of opaque minerals, chlorite and quartz.
Analytical Results
Whole-rock geochemistry
Representative rocks from the Bornaward (Taknar) granitoid were analyzed for major and trace elements (Table 1). Chemical composition of Bornaward (Taknar) granitoid rocks were plotted in TAS diagram [14], and they fall into the fields of granite and granodiorite (Fig. 7). Villaseca [15] divided per-aluminous granites into four groups (Fig. 8): (1) Highly per-aluminous granitoids (H-P). These rocks have the highest A = (Al/(K + Na +2Ca). They are characterized by having Al-rich minerals such as muscovite, garnet (almandine-spessartine series), cordierite, and sillimanite. They usually contain abundant restitic enclaves. (2) Moderately per-aluminous granitoids (M-P). These are biotite bearing. Accessory phases are cordierite and garnet (almandine-spessartine series). (3) Low per-aluminous granitoids (L-P). These rocks could be evolved from I-type or low-ASI-type granitoids. They may contain amphibole. Enclaves are mainly mafic granular types. (4) Highly felsic per-aluminous granitoids (F-P).

Bornaward (Taknar) granite and granodiorite plot in the field of highly per-aluminous (m-P) and metaluminous granitoids (h-P) (Fig. 8). The granitic rocks are medium- to high-potassic (Fig. 9).

Chemical composition of Bornaward (Taknar) granitoid rocks are plotted in Ga/Al- vs. trace element discrimination diagrams [18]. They plot in the field of I & S- type granites (Fig. 10).

<table>
<thead>
<tr>
<th>Sample Location</th>
<th>TK2-8</th>
<th>TK1-5</th>
<th>TK2-18</th>
<th>TK1-48</th>
<th>TK2-35</th>
<th>TK2-29</th>
<th>TK2-15</th>
<th>TK2-18</th>
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<tr>
<td>SiO₂</td>
<td>71.37</td>
<td>72.68</td>
<td>76</td>
<td>69.77</td>
<td>71.55</td>
<td>71.61</td>
<td>71.48</td>
<td>70.8</td>
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<td>TiO₂</td>
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<td>0.26</td>
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<td>0.16</td>
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<tr>
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<td>1.08</td>
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<td>2.49</td>
<td>2.12</td>
<td>3.09</td>
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<td>0.03</td>
<td>0.04</td>
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<td>2.66</td>
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<td>3.36</td>
<td>3.71</td>
<td>4.99</td>
<td>3.49</td>
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<td>0.06</td>
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<td>0.06</td>
<td>0.04</td>
<td>0.08</td>
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<td>127.3</td>
<td>124.3</td>
<td>130.4</td>
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<td>40</td>
<td>36.8</td>
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<td>110.9</td>
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<td>1656</td>
<td>1755</td>
<td>1311</td>
<td>1299</td>
<td>1974</td>
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<td>137.7</td>
<td>167.8</td>
<td>168.9</td>
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Figure 7. Composition of Bornaward (Taknar) granitoid plotted on the TAS diagram [14].
Figure 8. Composition of Bornaward (Taknar) granitoid plotted on Villaseca [15] diagram. Bornaward (Taknar) granite and granodiorite plot in the fields of moderately peraluminous (M-P) and high peraluminous. (L-P = low peraluminous; M-P = moderately peraluminous; H-P = highly peraluminous; F-P = felsic peraluminous).

Figure 9. K$_2$O vs. SiO$_2$ variation diagram [16] with boundaries by [17] for high-K, medium-K, and low-K magma series.

Figure 10. Bornaward granitoid are plotted in the field of I & S type [18].
Compositions of granitic rocks from the Bornaward plotted on the CaO/Na2O–Al2O3/TiO2 diagram. The field of S-granites and its boundaries are given after [19]. Bornaward granitoid plot in the field of S–type granite (Fig. 11). Representative rock samples from Bornaward granite and granodiorite were analyzed for rare earth and trace elements using ICP-MS at the ACME Lab (Canada) (Table 2). REE from Dehnow S-Type Paleo-Tethys granitoids (Mashhad) and Najmabad granitoids (Ghonabad) (both S and I–type’s granitoids) are presented in Table 2 [20, 21]. Lower continental crust-normalized trace-element diagrams indicates that both granite and granodiorite are relatively enriched in LILE = Rb, K and depleted in Cs, Ba, Nb, Ta, Sr (Fig. 12). In comparison with granodiorite (S-type) from Dehnow and Najmabad, Bornaward granitoids show very strong depletion in Sr and Ti (Fig. 12).

Chondrite-normalized rare earth element (REE) plots indicate minor enrichments of light relative to heavy REE, with (La/Yb)N between 3.5-5.6 (Fig. 13) and high total REE = 193-252 with strong negative anomaly of Eu (Fig. 13). Dehnow and Najmabad granodiorite have (La/Yb)N = 7 to 22 and no, or only small, negative Eu anomalies (Eu/Eu* = 0.55 to 1.1) (Fig. 13).

**Figure 11.** Compositions of granitic rocks from the Bornaward plotted on the CaO/Na2O–Al2O3/TiO2 diagram. The field of S-granites and its boundaries types are given after [19]. Bornaward granitoids are plotted in the field of S–type granite.

**Table 2.** REE elements analysis of Bornaward, Dehnow and Najmabad intrusive rocks [20, 21].

<table>
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<tr>
<th>Location</th>
<th>Sample</th>
<th>La</th>
<th>Ce</th>
<th>Pr</th>
<th>Nd</th>
<th>Sm</th>
<th>Eu</th>
<th>Gd</th>
<th>Tb</th>
<th>Dy</th>
<th>Ho</th>
<th>Er</th>
<th>Tm</th>
<th>Yb</th>
<th>Lu</th>
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<td>Bornaward</td>
<td>Tk2-18 Granite</td>
<td>42.1</td>
<td>82.6</td>
<td>8.7</td>
<td>33.3</td>
<td>7.18</td>
<td>0.64</td>
<td>7.59</td>
<td>1.58</td>
<td>10.04</td>
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<td>Tk2-8 Granite</td>
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<td>Granodiorite</td>
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<td>0.6</td>
<td>0.07</td>
<td>0.55</td>
<td>0.08</td>
</tr>
<tr>
<td>Najmabad</td>
<td>KP-2 Monzonite</td>
<td>16.5</td>
<td>37.8</td>
<td>4.57</td>
<td>18.1</td>
<td>2.98</td>
<td>0.85</td>
<td>2.2</td>
<td>0.28</td>
<td>1.33</td>
<td>0.21</td>
<td>0.56</td>
<td>0.06</td>
<td>0.5</td>
<td>0.06</td>
</tr>
<tr>
<td>Najmabad</td>
<td>KP-25 Monzonite</td>
<td>16.9</td>
<td>37.1</td>
<td>4.36</td>
<td>16.8</td>
<td>2.9</td>
<td>0.82</td>
<td>2.04</td>
<td>0.27</td>
<td>1.33</td>
<td>0.21</td>
<td>0.63</td>
<td>0.07</td>
<td>0.53</td>
<td>0.08</td>
</tr>
</tbody>
</table>
Figure 12. Spider diagram for Bornaward, Dehnow, and Najmabad intrusive. Data for Lower continental crust are from [22].

Figure 13. Chondrite-normalized REE distribution [23] for Bornaward granite and granodiorite.

Magnetic susceptibility
Granitic rocks were classified into the magnetite series and the ilmenite series by Ishihara (1977). Ishihara recognized that in Japan there is a distinct spatial distribution of granitic rocks that contain magnetite coexisting with ilmenite and those that contain ilmenite as the only Fe-Ti oxide. He recognized that the magnetite-series granitoids are relatively oxidized, whereas the ilmenite-series granitoids are relatively reduced. Granites showing a magnetic susceptibility value of $>3.0 \times 10^{-3}$ (SI units) are classified as belonging to the magnetite series [24]. Magnetic susceptibility ($<3.0 \times 10^{-3}$ (SI units) and mineralogical composition of the studied granite and granodiorite indicate that they belong to the ilmenite series.

U-Th-Pb Zircon age dating
The age of Taknar Formation and the granitoid rocks (Bornaward) are becoming an interesting subject. Based on field observation, Muller and Walter [1] suggested that Taknar Formation is Precambrian in age. Recent Palynological studies show that the age of Taknar Formation is Ordovician [4]. Using the Rb–Sr whole-rock and biotite isotopic methods by Soltani [5] on granitoid rocks (Bornaward granite) gave values of 154 to 111 Ma. The initial $(\frac{87}{86}\text{Sr}) > 0.76$ for the granitoid based on Soltani [5] is very high and unusual. It became interested to find out the accurate age of Taknar Formation and the granitoids. We used more accurate U-Pb in zircon age dating.

Two rock samples, which were analyzed for Sr and Nd isotopes (granodiorite-granite), were chosen for U-Pb-Th dating of Zircon. The U-Th-Pb dating of Zircon was done at Department of Geosciences University of Arizona Tucson, (USA).

The results of U-Th-Pb zircon analysis of the two samples No TAK2-8 and TAK1-5 are presented in the Table (3). The results of calculation of isotopic (sample no TAK2-8, granite) age is presented in the TuffZirc diagram (Fig. 12). The zircon sample TAK2-8 (granite) (Fig. 12) has 12 analyzed points giving the mean
age value (weighted mean) of 552.69 ± 10.89 Ma (error in the 2 sigma level). The results of calculation of isotopic (sample TAK1-5, granodiorite) age are presented in the TuffZirc graphics (Fig. 13). Based on 32 analyzed points the mean age value (weighted mean) for granodiorite is 538.22 – 1.82, ± 4.28 Ma (error in the 2 sigma level).

Table 3. Results of U-Pb-Th laser-ablation multicollector ICP mass spectrometry analysis of zircon from Bornaward granite and granodiorite.

<table>
<thead>
<tr>
<th>Analysis</th>
<th>U (ppm)</th>
<th>206Pb/204Pb</th>
<th>U/Th</th>
<th>206Pb/207Pb ± (%)</th>
<th>207Pb/235U ± (%)</th>
<th>206Pb/238U ± (%)</th>
<th>Age (Ma) ± (Ma)</th>
</tr>
</thead>
<tbody>
<tr>
<td>TAK1-5-1C</td>
<td>432</td>
<td>45846</td>
<td>3.4</td>
<td>14.9263 ± 0.8</td>
<td>1.2464 ± 0.1</td>
<td>0.1349 ± 0.1</td>
<td>815.9 ± 11.1</td>
</tr>
<tr>
<td>TAK1-5-1C</td>
<td>461</td>
<td>40369</td>
<td>2.1</td>
<td>16.9605 ± 1.4</td>
<td>0.7346 ± 0.4</td>
<td>0.0940 ± 0.3</td>
<td>557.7 ± 1.8</td>
</tr>
<tr>
<td>TAK1-5-3T</td>
<td>305</td>
<td>34228</td>
<td>2.3</td>
<td>17.3272 ± 2.8</td>
<td>0.6938 ± 2.9</td>
<td>0.0872 ± 0.7</td>
<td>538.9 ± 3.8</td>
</tr>
<tr>
<td>TAK1-5-4T</td>
<td>428</td>
<td>21930</td>
<td>2.8</td>
<td>17.0480 ± 1.5</td>
<td>0.6695 ± 1.7</td>
<td>0.0828 ± 0.7</td>
<td>512.7 ± 3.6</td>
</tr>
<tr>
<td>TAK1-5-5C</td>
<td>564</td>
<td>99521</td>
<td>1.1</td>
<td>13.9042 ± 0.6</td>
<td>1.6024 ± 4.0</td>
<td>0.1616 ± 4.0</td>
<td>983.7 ± 11.9</td>
</tr>
<tr>
<td>TAK1-5-6T</td>
<td>1071</td>
<td>94482</td>
<td>2.9</td>
<td>16.4676 ± 0.9</td>
<td>0.8479 ± 1.8</td>
<td>0.1013 ± 1.6</td>
<td>621.9 ± 9.6</td>
</tr>
<tr>
<td>TAK1-5-7C</td>
<td>448</td>
<td>30356</td>
<td>3.1</td>
<td>17.0575 ± 1.2</td>
<td>0.6981 ± 1.7</td>
<td>0.0864 ± 1.2</td>
<td>534.0 ± 6.0</td>
</tr>
<tr>
<td>TAK1-5-8C</td>
<td>360</td>
<td>33821</td>
<td>3.7</td>
<td>16.8601 ± 1.4</td>
<td>0.7605 ± 4.9</td>
<td>0.0930 ± 4.7</td>
<td>573.2 ± 25.7</td>
</tr>
<tr>
<td>TAK1-5-9C</td>
<td>1202</td>
<td>112969</td>
<td>2.0</td>
<td>17.0753 ± 0.4</td>
<td>0.6973 ± 1.4</td>
<td>0.0864 ± 1.3</td>
<td>533.9 ± 6.7</td>
</tr>
<tr>
<td>TAK1-5-10C</td>
<td>529</td>
<td>35368</td>
<td>3.8</td>
<td>16.9637 ± 0.7</td>
<td>0.6901 ± 1.4</td>
<td>0.0849 ± 1.1</td>
<td>525.3 ± 5.7</td>
</tr>
<tr>
<td>TAK1-5-11C</td>
<td>290</td>
<td>21759</td>
<td>4.1</td>
<td>17.0125 ± 3.2</td>
<td>0.6766 ± 3.4</td>
<td>0.0835 ± 1.3</td>
<td>516.9 ± 6.6</td>
</tr>
<tr>
<td>TAK1-5-12T</td>
<td>329</td>
<td>33909</td>
<td>3.5</td>
<td>17.0886 ± 1.4</td>
<td>0.7222 ± 1.7</td>
<td>0.0895 ± 1.0</td>
<td>552.6 ± 5.2</td>
</tr>
<tr>
<td>TAK1-5-13T</td>
<td>446</td>
<td>42334</td>
<td>2.6</td>
<td>17.0743 ± 1.3</td>
<td>0.7181 ± 2.0</td>
<td>0.0889 ± 1.5</td>
<td>549.2 ± 7.8</td>
</tr>
<tr>
<td>TAK1-5-14T</td>
<td>938</td>
<td>61488</td>
<td>2.7</td>
<td>16.8978 ± 0.7</td>
<td>0.7317 ± 2.5</td>
<td>0.0897 ± 2.4</td>
<td>553.6 ± 12.7</td>
</tr>
<tr>
<td>TAK1-5-16T</td>
<td>271</td>
<td>15615</td>
<td>1.7</td>
<td>16.6030 ± 4.0</td>
<td>0.7184 ± 4.9</td>
<td>0.0865 ± 2.8</td>
<td>534.8 ± 14.2</td>
</tr>
<tr>
<td>TAK1-5-17T</td>
<td>542</td>
<td>40302</td>
<td>1.6</td>
<td>16.8853 ± 1.2</td>
<td>0.6911 ± 2.5</td>
<td>0.0846 ± 2.2</td>
<td>523.7 ± 11.3</td>
</tr>
<tr>
<td>TAK1-5-18C</td>
<td>510</td>
<td>34756</td>
<td>1.3</td>
<td>16.9926 ± 1.3</td>
<td>0.7133 ± 1.5</td>
<td>0.0879 ± 0.8</td>
<td>543.2 ± 4.4</td>
</tr>
<tr>
<td>TAK1-5-19T</td>
<td>376</td>
<td>30795</td>
<td>2.1</td>
<td>17.0282 ± 2.2</td>
<td>0.7159 ± 3.9</td>
<td>0.0884 ± 3.2</td>
<td>546.1 ± 16.8</td>
</tr>
<tr>
<td>TAK1-5-20T</td>
<td>258</td>
<td>27304</td>
<td>2.1</td>
<td>17.3455 ± 1.4</td>
<td>0.6987 ± 1.7</td>
<td>0.0879 ± 0.9</td>
<td>543.1 ± 4.7</td>
</tr>
<tr>
<td>TAK1-5-21T</td>
<td>400</td>
<td>26329</td>
<td>3.1</td>
<td>16.7358 ± 2.8</td>
<td>0.7597 ± 3.5</td>
<td>0.0922 ± 2.1</td>
<td>568.6 ± 11.2</td>
</tr>
<tr>
<td>TAK1-5-22T</td>
<td>313</td>
<td>23018</td>
<td>1.6</td>
<td>17.0354 ± 2.3</td>
<td>0.7057 ± 4.8</td>
<td>0.0872 ± 4.2</td>
<td>538.9 ± 21.9</td>
</tr>
<tr>
<td>TAK1-5-23T</td>
<td>738</td>
<td>67172</td>
<td>1.4</td>
<td>17.0728 ± 0.9</td>
<td>0.7057 ± 1.4</td>
<td>0.0874 ± 1.1</td>
<td>540.0 ± 5.6</td>
</tr>
<tr>
<td>TAK1-5-24T</td>
<td>481</td>
<td>53142</td>
<td>2.4</td>
<td>17.0398 ± 0.9</td>
<td>0.7065 ± 1.6</td>
<td>0.0873 ± 1.3</td>
<td>539.7 ± 6.9</td>
</tr>
<tr>
<td>TAK1-5-25T</td>
<td>227</td>
<td>19898</td>
<td>2.8</td>
<td>16.8117 ± 2.1</td>
<td>0.7082 ± 2.8</td>
<td>0.0864 ± 1.9</td>
<td>533.9 ± 9.7</td>
</tr>
<tr>
<td>TAK1-5-27T</td>
<td>335</td>
<td>11428</td>
<td>3.3</td>
<td>16.6330 ± 5.4</td>
<td>0.7209 ± 5.5</td>
<td>0.0870 ± 0.7</td>
<td>537.6 ± 3.4</td>
</tr>
<tr>
<td>TAK1-5-28T</td>
<td>573</td>
<td>27325</td>
<td>1.7</td>
<td>16.9589 ± 1.3</td>
<td>0.7047 ± 2.4</td>
<td>0.0867 ± 2.0</td>
<td>535.8 ± 10.3</td>
</tr>
<tr>
<td>TAK1-5-29T</td>
<td>258</td>
<td>31373</td>
<td>1.9</td>
<td>16.3340 ± 1.2</td>
<td>0.8920 ± 2.4</td>
<td>0.1057 ± 2.1</td>
<td>647.6 ± 12.6</td>
</tr>
<tr>
<td>TAK1-5-30T</td>
<td>613</td>
<td>18086</td>
<td>1.7</td>
<td>16.8825 ± 1.6</td>
<td>0.7123 ± 2.7</td>
<td>0.0872 ± 2.2</td>
<td>539.1 ± 11.3</td>
</tr>
<tr>
<td>TAK1-5-31T</td>
<td>412</td>
<td>35728</td>
<td>2.9</td>
<td>16.9675 ± 2.0</td>
<td>0.6992 ± 2.9</td>
<td>0.0860 ± 2.0</td>
<td>532.1 ± 10.3</td>
</tr>
<tr>
<td>TAK1-5-32T</td>
<td>421</td>
<td>23796</td>
<td>4.0</td>
<td>16.7739 ± 1.5</td>
<td>0.6941 ± 2.4</td>
<td>0.0844 ± 1.9</td>
<td>522.5 ± 9.7</td>
</tr>
</tbody>
</table>
Sr–Nd Isotopes

Sr and Nd isotope data for representative samples are given in Table (4). They have a range in initial \(^{87}\text{Sr}/^{86}\text{Sr}\) and \(^{143}\text{Nd}/^{144}\text{Nd}\) from 0.713566 to 0.716888 and 0.511791 to 0.511842, respectively, when recalculated to an age of 553 and 538 Ma, they are consistent with the new radiometric results (Table 4). Initial \(\varepsilon\text{Nd}\) isotope values for granite and granodiorite range from -2.62 to -2.01 (Table 4).

Soltani [5] analyzed several plutonic rocks between Torbat-Hydarieh, Drouneh and Sabzevar for Rb-Sr. The result of his study only for two localities, Bornaward (Taknar) and Kashmar are presented in Table (5). Based on field data and Rb-Sr age, the Kashmar granitoid formed in Tertiary. The initial \(^{87}\text{Sr}/^{86}\text{Sr}\) is between 0.7048 to 0.7056 and the age is 42 Ma (Table 5). The initial \(^{87}\text{Sr}/^{86}\text{Sr}\) for Bornaward granitoid is between 0.71637 to 0.7500 and the age is 152.8 to 111.8 Ma (Table 5). The age which was calculated based on Rb-Sr in whole rock and biotite is incorrect (Table 5). Since the age was incorrect therefore the initial \(^{87}\text{Sr}/^{86}\text{Sr}\) is incorrect too.
Table 4a. Rb-Sr isotopic composition of Bornaward granite and granodiorite

<table>
<thead>
<tr>
<th>Sample</th>
<th>Rb (ppm)</th>
<th>Sr (ppm)</th>
<th>(^{87}\text{Rb}/^{86}\text{Sr})</th>
<th>(^{87}\text{Sr}/^{86}\text{Sr})_m (2σ)</th>
<th>R0(Sr)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Granodiorite TK1-5</td>
<td>111</td>
<td>50</td>
<td>6.0167</td>
<td>0.766090 (9)</td>
<td>0.716888</td>
</tr>
<tr>
<td>Granite TK2-8</td>
<td>41</td>
<td>79</td>
<td>1.4999</td>
<td>0.725390 (1)</td>
<td>0.713566</td>
</tr>
</tbody>
</table>

Table 4b. Sm-Nd isotopic composition of Bornaward granite and granodiorite

<table>
<thead>
<tr>
<th>Sample</th>
<th>Sm (ppm)</th>
<th>Nd (ppm)</th>
<th>(^{147}\text{Sm}/^{144}\text{Nd})</th>
<th>(^{143}\text{Nd}/^{144}\text{Nd})_m (2σ)</th>
<th>R0(Nd)</th>
<th>εNd</th>
<th>TDM</th>
</tr>
</thead>
<tbody>
<tr>
<td>Granodiorite TK1-5</td>
<td>6.18</td>
<td>27</td>
<td>0.1385</td>
<td>0.512330 (12)</td>
<td>0.511842</td>
<td>-2.01</td>
<td>1.44</td>
</tr>
<tr>
<td>Granite TK2-8</td>
<td>7.7</td>
<td>36</td>
<td>0.1294</td>
<td>0.512260 (09)</td>
<td>0.511791</td>
<td>-2.62</td>
<td>1.41</td>
</tr>
</tbody>
</table>

m= measured. Errors are reported as 2σ (95% confidence limit).

R0(Sr) is the initial ratio of \(^{87}\text{Sr}/^{86}\text{Sr}\) for each sample, calculated using \(^{87}\text{Rb}/^{86}\text{Sr}\) and \(^{87}\text{Sr}/^{86}\text{Sr}\)_m and an age of 538 Ma (granodiorite) and 553 Ma (granite) (age based on zircon).

R0(Nd) is the initial ratio of \(^{143}\text{Nd}/^{144}\text{Nd}\) for each sample, calculated using \(^{147}\text{Sm}/^{144}\text{Nd}\) and \(^{143}\text{Nd}/^{144}\text{Nd}\)_m and an age of 538 Ma (granodiorite) and 553 Ma (granite) (age based on zircon).

BNdI, initial BNd value.

Table 5. Rb-Sr and Age of Bornaward (Taknar) and Kashmar granite [5]

<table>
<thead>
<tr>
<th>Sample</th>
<th>Rock</th>
<th>Rb (ppm)</th>
<th>Sr (ppm)</th>
<th>(^{87}\text{Rb}/^{86}\text{Sr})</th>
<th>(^{87}\text{Sr}/^{86}\text{Sr})_m (2σ)</th>
<th>(^{87}\text{Sr}/^{86}\text{Sr})</th>
<th>Age Ma</th>
</tr>
</thead>
<tbody>
<tr>
<td>R15947</td>
<td>Granodiorite</td>
<td>70</td>
<td>109</td>
<td>1.904</td>
<td>0.72050</td>
<td>0.71637</td>
<td>152.8 ±1.3</td>
</tr>
<tr>
<td></td>
<td>Biotite</td>
<td>622</td>
<td>148.8</td>
<td>124.633</td>
<td>0.98700</td>
<td></td>
<td></td>
</tr>
<tr>
<td>R15946</td>
<td>Granodiorite</td>
<td>129</td>
<td>145</td>
<td>2.641</td>
<td>0.72700</td>
<td>0.72153</td>
<td>145.6±1.3</td>
</tr>
<tr>
<td></td>
<td>Biotite</td>
<td>635</td>
<td>13</td>
<td>148.788</td>
<td>1.02937</td>
<td></td>
<td></td>
</tr>
<tr>
<td>R15938</td>
<td>granite</td>
<td>95</td>
<td>39</td>
<td>7.246</td>
<td>0.75253</td>
<td>0.73978</td>
<td>123.8±1</td>
</tr>
<tr>
<td></td>
<td>biotite</td>
<td>947</td>
<td>11</td>
<td>258.519</td>
<td>1.19453</td>
<td></td>
<td></td>
</tr>
<tr>
<td>R15941</td>
<td>granite</td>
<td>128</td>
<td>39</td>
<td>9.775</td>
<td>0.76562</td>
<td>0.75008</td>
<td>111.8±1.1</td>
</tr>
<tr>
<td></td>
<td>biotite</td>
<td>715</td>
<td>16.17</td>
<td>130.959</td>
<td>0.95823</td>
<td></td>
<td></td>
</tr>
</tbody>
</table>

<table>
<thead>
<tr>
<th>Sample</th>
<th>Rock</th>
<th>Rb (ppm)</th>
<th>Sr (ppm)</th>
<th>(^{87}\text{Rb}/^{86}\text{Sr})</th>
<th>(^{87}\text{Sr}/^{86}\text{Sr})_m (2σ)</th>
<th>(^{87}\text{Sr}/^{86}\text{Sr})</th>
<th>Age Ma</th>
</tr>
</thead>
<tbody>
<tr>
<td>R15915</td>
<td>Granodiorite</td>
<td>62</td>
<td>273</td>
<td>0.672</td>
<td>0.70610</td>
<td>0.70569</td>
<td>42.8±0.5</td>
</tr>
<tr>
<td></td>
<td>Biotite</td>
<td>346</td>
<td>23.3</td>
<td>43.115</td>
<td>0.73193</td>
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<td></td>
</tr>
<tr>
<td>R15910</td>
<td>Granite</td>
<td>103</td>
<td>315</td>
<td>0.968</td>
<td>0.70574</td>
<td>0.70516</td>
<td>42.4±0.4</td>
</tr>
<tr>
<td></td>
<td>Biotite</td>
<td>440</td>
<td>7.9</td>
<td>162.759</td>
<td>0.80313</td>
<td></td>
<td></td>
</tr>
<tr>
<td>R15918</td>
<td>Granite</td>
<td>88</td>
<td>288</td>
<td>1.382</td>
<td>0.70615</td>
<td>0.70532</td>
<td>42.5±0.5</td>
</tr>
<tr>
<td></td>
<td>biotite</td>
<td>397</td>
<td>14.4</td>
<td>77.998</td>
<td>0.75235</td>
<td></td>
<td></td>
</tr>
<tr>
<td>R15900</td>
<td>Alkali feldspar granite</td>
<td>200</td>
<td>50</td>
<td>11.83</td>
<td>0.71209</td>
<td>0.70478</td>
<td>43.5±0.4</td>
</tr>
<tr>
<td></td>
<td>biotite</td>
<td>689.8</td>
<td>4.53</td>
<td>16.99</td>
<td>0.71504</td>
<td></td>
<td></td>
</tr>
</tbody>
</table>

Discussion section

Source of Magma

Based on mineralogy and low values of magnetic susceptibility [(5 to 11) \times 10^{-5} \text{SI}], both granite and granodiorite are classified as belonging to the ilmenite-series of reduced S-type granitoids. Chemically they are per-aluminous, high-K calc-alkaline with relative enrichment in LILE, Rb, K and depletion in Sr, Ba, Nb, Ti, Ta, Y, Yb. Chondrite-normalized rare earth element (REE) plots indicate minor enrichments of light relative to heavy REE, with (La/Yb)_N between 3.5-5.6. They show high total REE = 193-252 with strong negative anomaly of Eu.

Initial \(^{87}\text{Sr}/^{86}\text{Sr}\), εNd and \(^{143}\text{Nd}/^{144}\text{Nd}\) of MORB basalts, Bornaward Granite-granodiorite and Dehnow diorite-granodiorite are plotted in Fig. 16. Initial \(^{87}\text{Sr}/^{86}\text{Sr}\) isotope values for Bornaward Granite-granodiorite range from 0.713566 to 0.716888, Dehnow diorite-granodiorite range from 0.707949 to 0.708589 and MORB is less than 0.7040. Initial \(^{143}\text{Nd}/^{144}\text{Nd}\) isotope values for Bornaward Granite-granodiorite range from 0.511842-0.511791, Dehnow diorite - granodiorite range from 0.512059-0.512019 and for MORB is between 0.5130 to 0.5135. Both Bornaword and Dehnow granitoid are originated from the continental crust. The Initial \(^{87}\text{Sr}/^{86}\text{Sr}\) isotope values for Bornaward Granite-granodiorite
(0.713566) is much higher than Dehnow granodiorite (0.708). This indicates the source of magma for Bornaward Granite-granodiorite was originated from the continental crust and is was more radiogenic.

Samarium-neodymium T_{DM}'s can provide an estimate for the time at which continental crust was extracted from a depleted mantle source. Therefore, T_{DM} ages may delineate crustal blocks of differing ages [25-27]. Granite and granodiorite of Bornaward yields a T_{DM} age of 1.4-1.41 Ga (Table 4b). This indicates that the granites and granodiorite being derived from partial melting of distinct basement source regions.

The Rb/Sr versus Rb/Ba discrimination diagram (Fig. 17) shows that Bornaward granitoid rocks plot next to metapelitic derived melt [19]. Dehydration-melting behavior of biotite in metapelitic rocks has been investigated by many workers, including [28-36]. The amount and composition of melts produced is dependent on different factors: composition of the source rocks, temperature, water, pressure and oxygen fugacity. In general, as the melt produced at low melting fractions are indistinguishable (mainly peraluminous), irrespective of the nature of the protolith only the amount of melt will be different. Even metaaluminous meta-igneous rocks such those reported by [37-41] are able to produce peraluminous felsic melts at low melting fractions and in water-deficient conditions. As the melting increases, primary component of the source rocks are progressively incorporated in the melt. The Ti and Mg content of biotite directly control the reaction temperature interval in biotite. As these elements increase in biotite, the reaction temperature interval will expand.

**Figure 16.** Plot of Initial $^{87}\text{Sr}/^{86}\text{Sr}$ versus Initial $^{143}\text{Nd}/^{144}\text{Nd}$ isotope values for Bornaward, Dehnow and Kashmar granitoid rocks.

**Figure 17.** Plot of Rb/Sr versus Rb/Ba, Bornaward granitoid plot in the field near politic rock but Dehnow granitoid plot in the field near meta-greywacke.
Late Neoproterozoic granitoid rocks in Iran
Results of study by Hassanzadeh [42] demonstrate that Late Neoproterozoic to Early Cambrian granitoids and granitic gneisses are present in all continental structural zones of Iran north of the Zagros, from the Sanandaj–Sirjan zone to the northern margin of the Alborz Mountains (Table 6). Therefore, the crystalline basement of Iran can be considered to be the approximate northern continuation of the Arabian platform [42]. Since definitive crust of comparable age is generally absent in cratonic Eurasia [43], the occurrences of granitoids with Neoproterozoic ages from the Zagros to the northern foothills of the Alborz point to a Gondwana affinity for the continental terrains composing Iran.

Granitoid rocks of Late Neoproterozoic age are exposed in Saghand region (central Iran) [44]. Two multi-grain and one single-grain zircon fractions from a Boneh-Shurow (Saghand region) quartz-diorite intrusion yielded concordant analyses with an average $^{207}\text{Pb}/^{206}\text{Pb}$ age of $547.6 \pm 2.5$ Ma [44] (Fig. 18). Figure (18) show the distribution of Late Neoproterozoic granitoid rocks in Iran.

Conclusions
Taknar zone is an exotic block bordered with two major faults, Great Kavir fault (Drouneh) to the south and Rivash fault in the north. Taknar Formation is divided into lower, middle and upper members. It consists of meta-tuff, meta-rhyolite, slate, chlorite schist and quartzite. A complex of intrusive rocks crop out at the center of Taknar zone. They are named as “Bornaward granite”. The intrusive rocks are composed of granite, granodiorite, monzonite and diorite. Granite and granodiorite are dominant. Based on field observation diorite is older than granite and granodiorite.

### Table 6. Thermal-ionization (TIMS) zircon U-Pb geochronology [42].

<table>
<thead>
<tr>
<th>Region</th>
<th>Area</th>
<th>Age Ma U-Pb Zircon</th>
</tr>
</thead>
<tbody>
<tr>
<td>Northwest Iran</td>
<td>Sarv-e Jahan granite</td>
<td>544±29 Ma</td>
</tr>
<tr>
<td></td>
<td></td>
<td>599±42 Ma</td>
</tr>
<tr>
<td></td>
<td>Doran granite</td>
<td>567±19 Ma</td>
</tr>
<tr>
<td>Takab-west Zanjan area</td>
<td>Moghanlou granite</td>
<td>548±27 Ma</td>
</tr>
<tr>
<td></td>
<td>Mahneshan granitoid complex</td>
<td>568±44 Ma</td>
</tr>
<tr>
<td>Northwestern Sanandaj–Sirjan zone</td>
<td>Sheikh Chupan granodiorite</td>
<td>551±25 Ma</td>
</tr>
<tr>
<td></td>
<td>Bubaktan</td>
<td>544±19 Ma</td>
</tr>
<tr>
<td>Golpaygan area</td>
<td>Muteh Au mine</td>
<td>578±22 Ma</td>
</tr>
<tr>
<td>Northeast central Iran</td>
<td>Torud granite</td>
<td>566±31 Ma</td>
</tr>
<tr>
<td></td>
<td>Khār Turan granite</td>
<td>554±40 Ma</td>
</tr>
<tr>
<td></td>
<td>Band-e Hezar Chāh granite</td>
<td>581±21 Ma</td>
</tr>
<tr>
<td>Northern Alborz</td>
<td>Lahijan granite</td>
<td>551±9 Ma</td>
</tr>
</tbody>
</table>
Both granite and granodiorite are classified as belonging to the ilmenite-series of reduced S-type granitoids. Chemically they are per-aluminous, high-K calc-alkaline with relative enrichment in LILE = Rb, K and depletion in Sr, Ba, Nb, Ti, Ta, Y, Yb. Chondrite-normalized rare earth element (REE) plots indicate minor enrichments of light relative to heavy REE, with \((\text{La/Yb})_N\) between 3.5-5.6. High total REE = 193-252 with strong negative anomaly of Eu.

The results of U-Pb zircon age of granodiorite is 552.69 ± 10.89 Ma and granite is 538.22 – 1.82, + 4.28 Ma (late Neoproterozoic time). They have a range in initial \(^{87}\text{Sr}/^{86}\text{Sr}\) and \(^{143}\text{Nd}/^{144}\text{Nd}\) from 0.713566 to 0.716888 and 0.511791 to 0.511842, respectively, when recalculated to an age of 553 and 538 Ma, consistent with the new radiometric results. Initial \(\varepsilon\text{Nd}\) isotope values for granite and granodiorite range from -2.62 to -2.01.

Granite and granodiorite of Bornaward yields a \(T_{\text{DM}}\) age of 1.4-1.41 Ga. This indicates that the granites and granodiorite being derived from partial melting of distinct basement source regions with very high initial \(^{87}\text{Sr}/^{86}\text{Sr}\).

Bornaward and other Late Neoproterozoic granitoid rocks which are exposed and are dated until today are located along the margin of central Iran plate. Knowing the petrogenesis and tectonic setting of all of these granitoids will help to understand the relationship between Gondwana and central Iran plate.

Acknowledgements
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References
types of granitoids, Eastern Najmabad, Ghonabad, Iran."., Geological Society of America (GSA) 2009.
[40] Patin’o-Douce A.E., Beard J.S., “Dehydration-melting of biotite gneiss and quartz